



I. Introduction

Campobello 7.5-minute quadrangle (CQ) is located on the Six Mile thrust sheet within the Tugaloo terrane of the Inner Piedmont. Geologic mapping of fifteen 7.5-minute quadrangles in the Western Inner Piedmont of South Carolina by Furman University and the South Carolina Geological Survey personnel (1995-2010) has investigated the polyphase fold deformation, faulting, and metamorphic history of the region. Our mapping traces these structures into CQ.



Photomicrograph (FOV 2mm) of hornblende-biotite-quartz-feldspar gneiss displaying hornblende (H), albite-twinned plagioclase (P), and titanite (T) (high relief). Crossed polars.



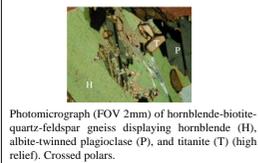
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II. Rock Type

Tallulah Falls Formation gneiss (Neoproterozoic) Interlayered biotite quartz feldspar gneiss, hornblende quartz feldspar gneiss, amphibolite, amphibole quartz feldspar gneiss, calc-silicate gneiss, and garnet biotite quartz feldspar gneiss. The different rock types are interlayered within CQ. The gneiss unit underlies ca. 95% of the quadrangle area.



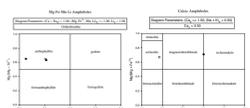
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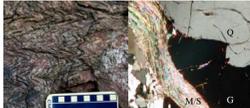


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Tallulah Falls Formation schist (Neoproterozoic) Interlayered schist and schistose muscovite-biotite gneiss.

The original stratigraphic order of metamorphic units in CQ is obscured by polyphase folding and faulting. Schists now lie structurally above the Tallulah Falls Formation gneiss. Schist and gneiss in the map unit have either biotite or muscovite or both present. Common lithologies include garnet (up to 5 mm)-mica schist; mica schist interlayered with pegmatite; fine-crystalline, limonitic-stained and beige-weathered, leucocratic, muscovite-sillimanite schist with sheared quartz lenses, 1-3mm thick and 0.5-2cm long; and a fine-crystalline, darker limonitic stained, muscovite schist.

With a decrease in modal mica and an increase in quartz and feldspar, schist grades into gneiss. The gneiss varieties include muscovite-biotite-garnet-quartz gneiss, with garnet comprising up to 40% of the rock composition and muscovite-biotite-quartz-plagioclase (oligoclase) gneiss.



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Holly Springs Hornblende Metagabbro A hornblende metagabbro body (3 x 0.5 km) is located between Lyman Lake and Holly Springs, SC. It is partly truncated and offset by east-northeast faults. Metagabbro is largely hornblende and plagioclase, the latter altering to epidote and scapolite in thin section. Some textures suggest the rock retains an original igneous texture. In float samples, zones of foliated amphibolite several cm wide with sharp contacts may represent original thin basaltic dikes in the gabbro.

The hornblende metagabbro body is surrounded by Tallulah Falls gneiss, but the contacts of the body are not exposed. Based on limited field data we tentatively suggest the hornblende metagabbro lies in the core of an F₂ synform (see cross section D-D'). If true, it is probably a fairly thin intrusive body. In addition, the hornblende metagabbro outcrop area between the North and Holly Springs faults may be a horst.



Photomicrograph (FOV 2mm) of recrystallized hornblende metagabbro. Hornblende (H) surrounds epidote (E) grains in the center of the image.

Silicified cataclastic rocks Microbreccia, cataclastite, and syntaxial quartz veins ("comb quartz") form narrow zones of outcrop, boulders, and resistant float. Discontinuous cataclastic rock bodies up to 1.3 km in length and trending N 60°-70°E adjacent to the Pax Mountain Fault have been traced across the entire Campobello quadrangle. On a ridge 1.6 km northeast of Liberty Church, two main zones of outcrops and boulders of microbreccia and comb quartz are 10m and 25m long. In general, linear cataclastic rock bodies are a few meters long and up to 3m high.



Microbreccia body 12 m wide of Pax Mountain fault zone and enclosing leucocratic biotite quartz-feldspar gneiss is exposed along Rt. 176 in Campobello, SC. View to northeast.



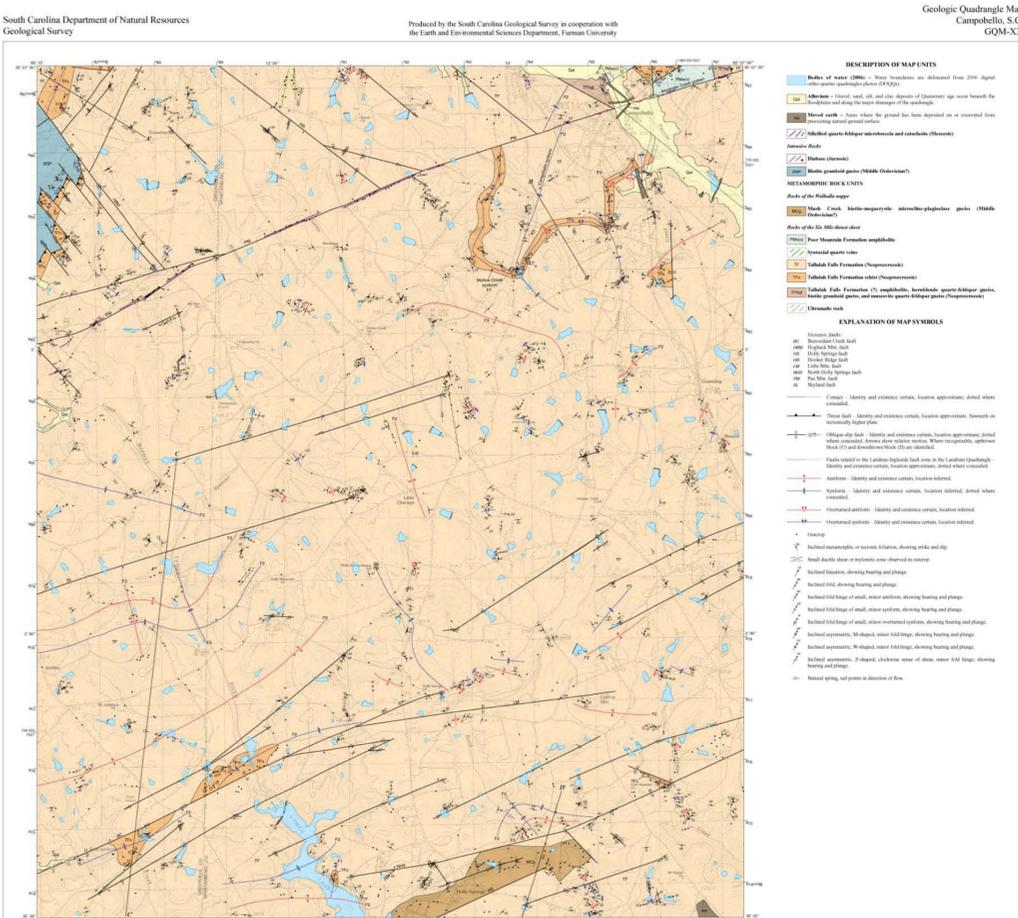
Thin, criss-crossing microbreccia and "comb quartz" veins (striking N40-80 E and N 18 W) are exposed over a 17m zone south of the 12m-wide Campobello microbreccia body. Pinkish enclosing rock is leucocratic biotite quartz-feldspar gneiss.



Continuous zone of microbreccia exposures and float (32 ft long) trends N45°E along Pax Mountain fault. Ridge 1.6 km northeast of Liberty Church.

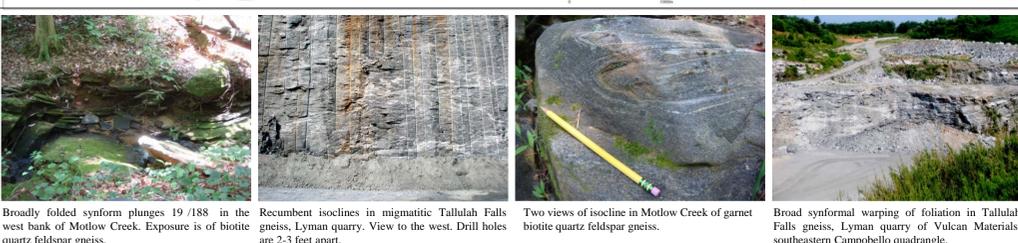
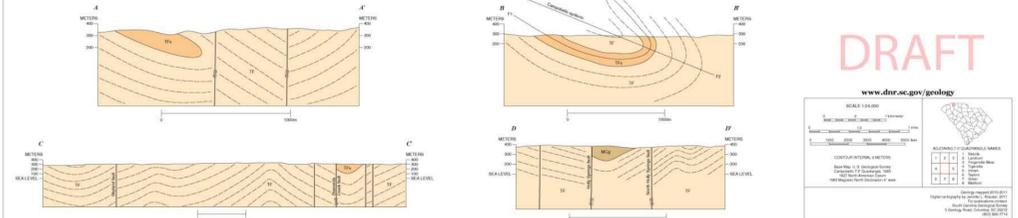


Photomicrograph (FOV 2mm) of microbreccia texture displaying clasts of quartz and feldspar of various sizes. Crossed polars.

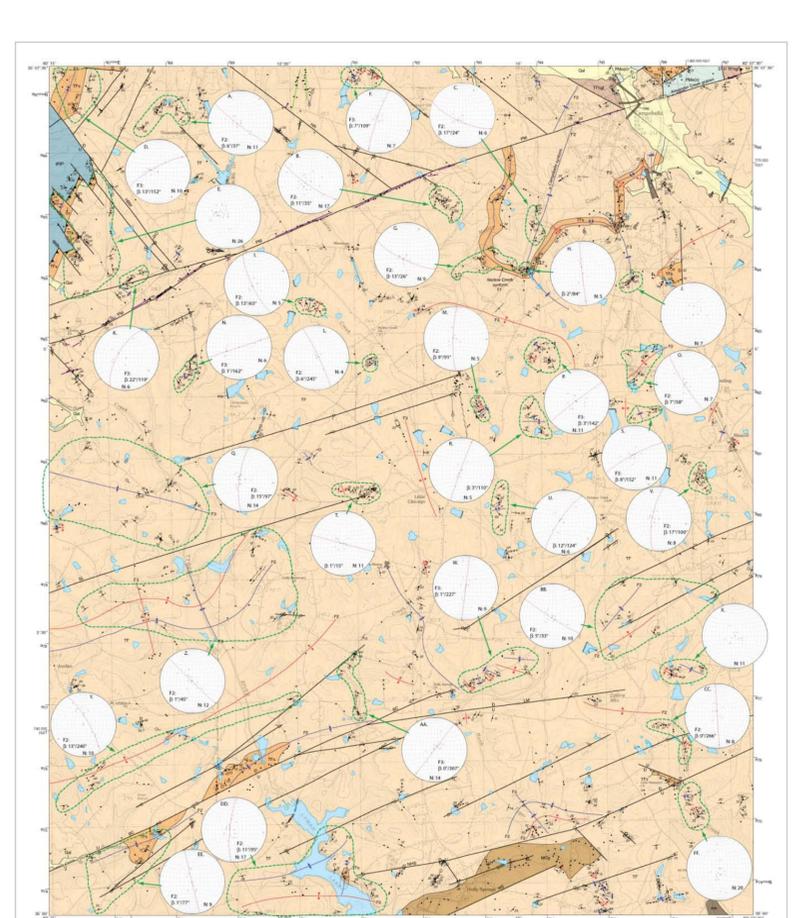


GEOLOGIC MAP OF THE CAMPOBELLO QUADRANGLE, GREENVILLE AND SPARTANBURG COUNTIES, SOUTH CAROLINA

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Photomicrograph (FOV 2mm) of microbreccia texture displaying clasts of quartz and feldspar of various sizes. Crossed polars.



SOUTHERN HEMISPHERE, EQUAL AREA STEREOPLOTS OF POLES TO FOLIATION SURFACES ARE USED TO DETERMINE STATISTICAL FOLD HINGE ORIENTATIONS (BETA) OF F₂ AND F₃ FOLDS

III. Cross Sections

A-A': In the northwest of CQ, Tallulah Falls schist lies structurally above Tallulah Falls gneiss in the nose of an overturned synform. The synform is truncated by a vertical fault. B-B': In the northeast of CQ and to the southeast of the Pax Mountain fault, schist forms a long, sinuous belt which has been deformed by F₁, F₂ and F₃ generation folds. The cross section shows part of this schist belt, which is structurally an overturned synform and consists of interlayered Tallulah Falls gneiss and schist. The earlier F₁ isoclinal fold was later refolded by F₂ folding. C-C': The Beaverdam Creek Fault, trending N55 E, in southeastern CQ truncates the schist unit, which occurs in the nose of a synform. The gneiss unit displays broad synforms and antiforms and is faulted by vertical faults. D-D': Hornblende Metagabbro body which is located in the southeast of CQ between Lyman Lake and Holly Springs, SC.

IV. Metamorphic History

Although metamorphism peaked in the sillimanite zone of the upper amphibolite facies, we observe a range of metamorphic conditions in the rocks of the CQ. Most notably, chlorite is developed in most of the amphibole quartz feldspar gneiss lithologies in which the amphibole is altered to chlorite. Microprobe data indicate that several amphiboles exist, which indicates a range of temperatures and pressures from the greenschist facies to the upper amphibolite facies. On the geologic map of Campobello quadrangle a trend appears for four of the locations of retrograded amphibole gneiss at N50 E and is a similar strike with nearby faults mapped in Campobello quadrangle. This similarity in trends suggests that the faults might have served as avenues for fluids, which then produced the retrograded mineralogy in localized regions.

V. Deformation: Faults and Folds

Folds in Campobello quadrangle F₁ - Isoclinal folds are found at all scales. The F₁ designation is used for the oldest folds we observed in CQ. Foliation and compositional layering are affected by F₁ folding. Hence it is unknown whether the compositional layering in gneiss is in fact axial planar to a fold set older than our F₁ set. Macroscopic F₁ folds were probably overturned (recumbent?) originally, with a northwest-vergente consistent with Griffin's nappe emplacement concepts for the Inner Piedmont (Wetmore, C.C., and Griffin, V.S., 1977; Griffin, V.S., 1971; Griffin, V.S., 1967). Aligned sillimanite needles may have formed during F₁ deformation. In rock exposures a transposition foliation has been observed.



Small isoclinal fold in biotite gneiss, Lyman quarry.

F₂ - Inclined to overturned macroscopic and mesoscopic F₂ folds deform earlier folds.



Overturned F₂ synform in interlayered biotite quartz-feldspar gneiss and amphibolite, Motlow Creek. Hammer lies on the gently dipping upright fold limb; to the right, layers abruptly bend and are overturned in the background at the narrow ledge crossing the creek. F₂ fold hinge plunges to the upper left at 30°/N59°E.

F₃ - Inclined to overturned macroscopic and mesoscopic F₃ folds deform earlier and F₂ folds.

F₄ and F₅ - Upright gentle folds. Not known from macroscopic map patterns



Pavement of biotite quartz feldspar gneiss along Meadow Creek, displaying a Type 1 interference fold pattern due to polyphase deformation. The white tape measures mark the axial traces of two gentle synforms which trend northwest and northeast.

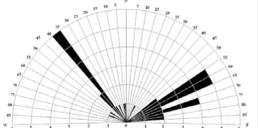
Polyphase folding in Campobello quadrangle

Campobello quadrangle has experienced five generations of folding resulting in polyphase deformation. The Motlow Creek F₁ isoclinal fold in the schist unit located in the northeast of CQ has been deformed by the macroscopic Campobello F₂ synform. Other F₂ folds are seen as tight chevron folds found in schist and have deformed F₁ aligned sillimanite needles, as evident from hooks that are visible in chevron folded schist. F₂ macroscopic inclined to overturned folds trend northwest and warp F₁ and F₃ folds. This can be seen clearly in the Campobello synform (F₂), which is cross folded by F₂ folds, F₃ and F₃ together create Type-2 interference folding. The final two generations of folding are mesoscopic F₄ and F₅, two generations of gentle folding that appear 90° to each other in stream pavements creating Type-1 interference folding.

Faults in Campobello quadrangle

The chronology of faulting in CQ is northeast (oldest), northerly, northwest, and east-northeast to easterly faults. Dominant fault sets consist of older and shorter N35 - 40 W striking faults and younger but longer N55 - 75 E striking faults.

Zones of microbreccia boulders lie along the Pax Mountain fault and trend N30 - 60 E, with individual sets trending N30 E, N40 - 50 E, N55 E and N60 E. In an exposure of microbreccia in a roadside ditch at GPS location 387668/3883602, near Morrow Road and less than 10m north of the trace of the Pax Mountain fault, extensional veins up to 1ft long were oriented N25 E 68 NW, N50 E 60 NW, and N55 E 82 NW. Northeast of this microbreccia exposure and still less than 10m from the Pax Mountain fault, the orientation of syntaxial veins was measured as N50 E, E-W, and N40 - 70 E. Two separate, north-trending syntaxial quartz veins (0.5-1 km long) lie 2 km northeast of Gowensville.



Rose Diagram showing fault strike frequencies for CQ



Pavement of biotite quartz feldspar gneiss along Meadow Creek, displaying a Type 1 interference fold pattern due to polyphase deformation. The white tape measures mark the axial traces of two gentle synforms which trend northwest and northeast.



Pavement of biotite quartz feldspar gneiss along Meadow Creek, displaying a Type 1 interference fold pattern due to polyphase deformation. The white tape measures mark the axial traces of two gentle synforms which trend northwest and northeast.

VI. Summary

The metamorphic peak is in the sillimanite zone of the upper amphibolite facies for the rocks in Campobello quadrangle. Localized retrograde fluids along faults trending N50 E metamorphosed rocks to greenschist facies, leading to chlorite replacement of amphiboles. Campobello quadrangle has experienced multiple deformational events resulting highly faulted and polyphase deformed geologic units. The fold chronology consists of five fold generations. F₁ mesoscopic and macroscopic isoclinal folds, F₂ inclined to overturned mesoscopic chevrons and macroscopic folds, F₃ gently inclined to overturned mesoscopic and macroscopic folds, and F₄ and F₅ gentle folds. The faulting chronology is (oldest-youngest): northeast, northerly, northwest, and east-northeast to northerly faults. Dominant fault trends are N35°-40°W and N55°-75°E and faults are assumed to be vertical.

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